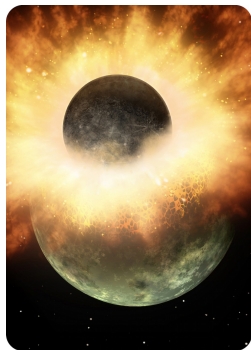


Constructing the Solar System: A Smashing Success

Constructing the Moon



Thomas M. Davison

Department of the Geophysical Sciences



THE UNIVERSITY OF
CHICAGO

Compton Lecture Series
Autumn 2012



Compton Lecture Series Schedule

- 1 10/06/12 A Star is Born
- 2 10/13/12 Making Planetesimals: The building blocks of planets
- 3 10/20/12 *Guest Lecturer: Mac Cathles*
- 4 10/27/12 Asteroids and Meteorites:
Our eyes in the early Solar System
- 5 11/03/12 Building the Planets
- 6 11/10/12 When Asteroids Collide
- 7 11/17/12 Making Things Hot: The thermal effects of collisions
11/24/12 No lecture: Thanksgiving weekend
- 8 12/01/12 Constructing the Moon
12/08/12 No lecture: Physics with a Bang!
- 9 12/15/12 Impact Earth: Chicxulub and other terrestrial impacts

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The Moon



Image courtesy of Gregory H. Revera/
Wikimedia Commons

Orbital distance	238854 miles
Average radius	1079 miles (0.27 × Earth)
Mass	7.34×10^{22} kg (0.012 × Earth)
Average density	3346 kg/m^3 (0.61 × Earth)

- Galileo 1609 showed Moon is rocky
 - Observed hills and what he thought were oceans
- Leads to our names for:
 - the bright highlands (**Terra**)
 - and dark basins (**Mare**)

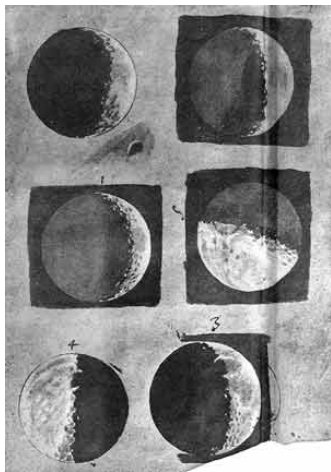


Image courtesy of The Galileo Project

Features on the Lunar surface

- **Maria** are large, dark, basaltic plain on the Moon
 - Formed by ancient volcanic eruptions
- **Terrae** are the lighter, anorthositic highlands
 - Older than the Mare

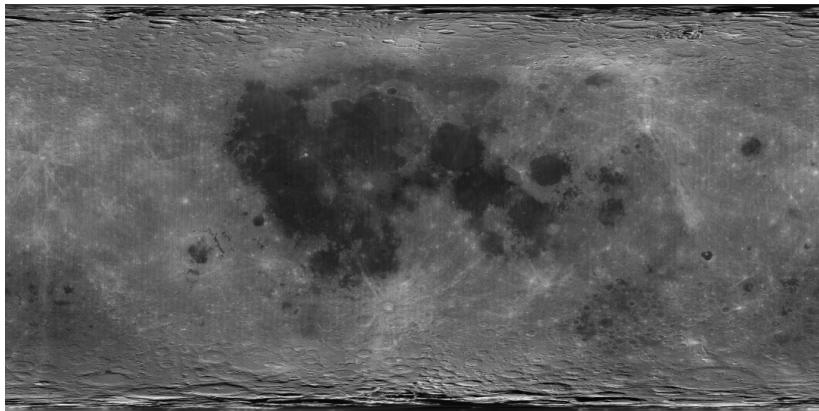
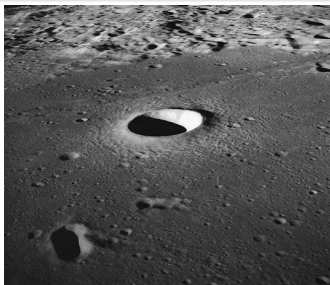
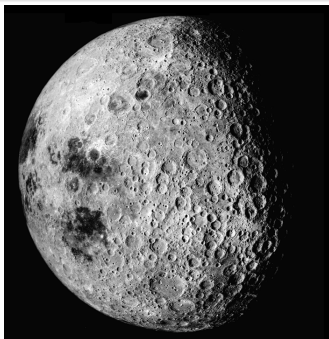


Image courtesy of the United States Geological Survey

Impact craters are a dominant surface feature



Images courtesy of NASA



- Craters cover the surface of the Moon
- Smooth areas are the younger Mare which often infill large craters

Craters still forming today

New Crater Discovered in LROC Image

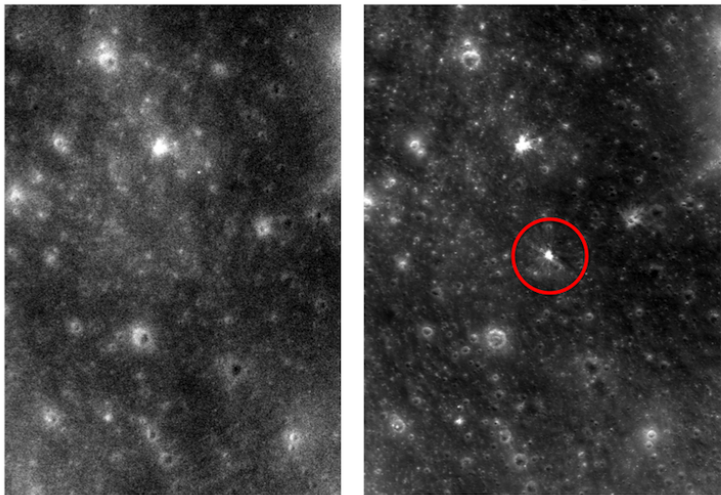


Image courtesy of NASA/GSFC/Arizona State University

Part 1:

Classical formation mechanisms



Image courtesy of NASA

- 1878: G.H. Darwin
 - Suggested a large chunk of material split from the Earth
 - Requires fast rotation of Earth
- Four years later: Osmond Fisher
 - Thought that Pacific basin represents the scar left behind

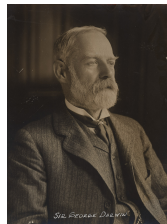
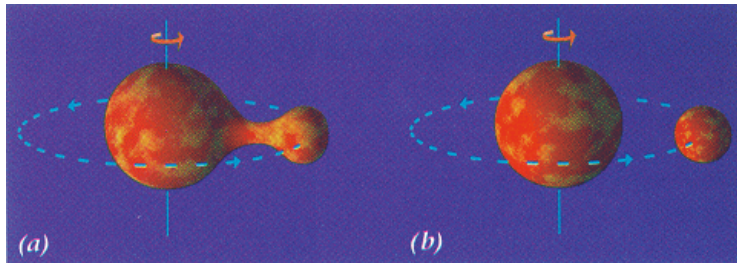


Image courtesy of J. Russell & Sons/
Wikimedia Commons



- 1909: Thomas Jefferson Jackson See
 - Moon formed as a planetesimal elsewhere in the Solar System
 - Later captured by Earth's gravity

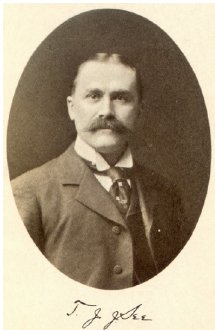


Image courtesy of William Larkin Webb/
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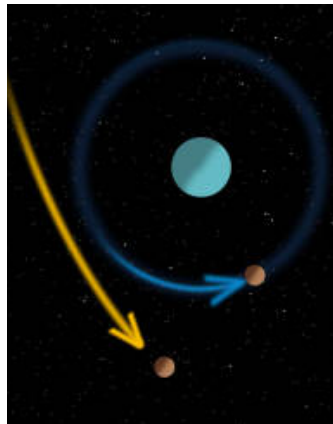
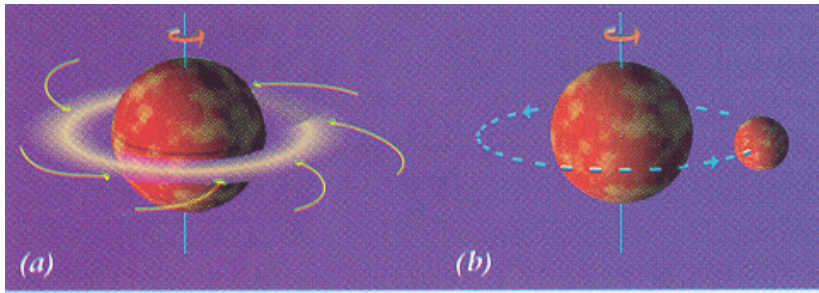


Image courtesy of Black Cat Studios

- 1873: Édouard Roche
 - Earth and Moon formed together
 - Formed from swarm of planetesimals as a **double planet**



Part 2:

Properties of our Moon



Image courtesy of NASA

Earth and the Moon: the view from Mars

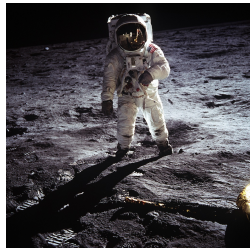
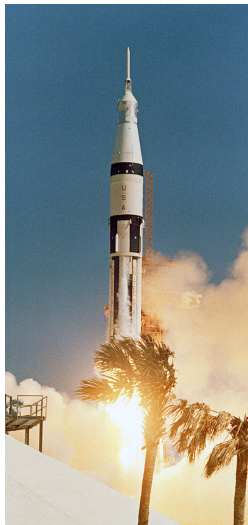
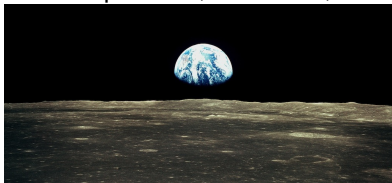


Image courtesy of NASA/JPL-Caltech/University of Arizona

- The Earth–Moon system has a high angular momentum
- The Moon has a greater mass fraction compared with other planet–satellite systems

We have samples of the Moon

- NASA's Apollo program in the 1960s to 1970s
- Collected samples from the Moon
 - Also have lunar meteorites
- Can measure ages, composition, structure, etc.



Images courtesy of NASA

The Moon formed late

- The Moon formed relatively late (given its size)
 - Radiometric age of lunar rocks from hafnium/tungsten isotopes
 - Moon formed > 30 million years after Solar System formation
 - Other objects this size formed in a few hundred thousand years

Harrison H. Schmitt collects lunar rocks during the Apollo 17 mission



Image courtesy of E.A. Cernan/NASA

The Moon has a small core

- Apollo astronauts left seismometers on the Moon
- Recorded moonquakes from 1969 to 1977
- Using moonquake data, the size of the lunar core is estimated to be ~ 330 km
 - $\sim 20\%$ of the Moon's radius
 - Compare to Earth's core ($\sim 50\%$ Earth's radius)
- Explains low density of Moon:
 - Moon is depleted in iron

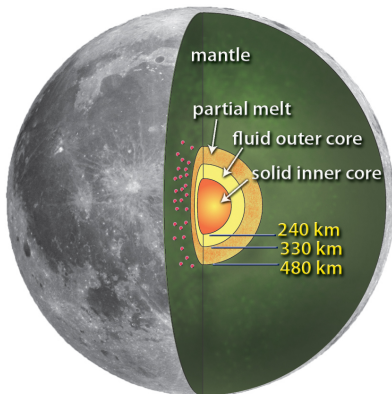


Image courtesy of NASA/MSFC/Renee Weber

The Moon had an early magma ocean

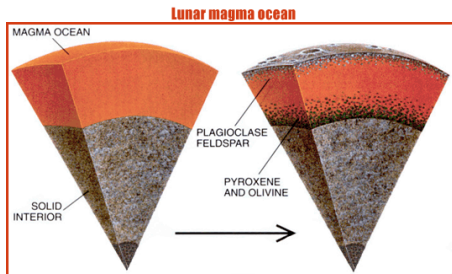


Image courtesy of G.J. Taylor
(*Scientific American/Planetary Science Research Discoveries*)

- The crust of the Moon is composed of a rock called **anorthosite**
- Anorthosite is a rock made up of low-density components
- This implies that the Moon had a magma ocean
 - The low-density components floated to the surface, and cooled as the crust

Summary of Constraints

A model of how the Moon formed should explain each of these

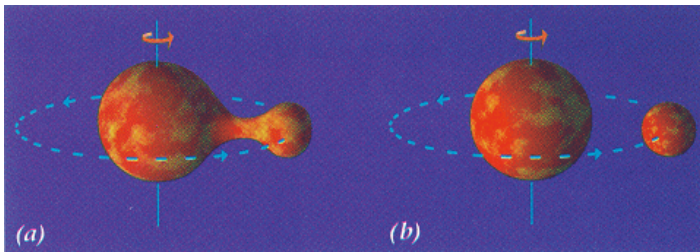
- 1 High angular momentum
- 2 Low density/small core
- 3 Late formation
- 4 Oldest rocks appear to come from magma ocean
 - Energetic start occurred after decay of short-lived radionuclides
- 5 Oxygen and titanium isotope data identical to Earth
 - Asteroids and planets differ
 - Common origin



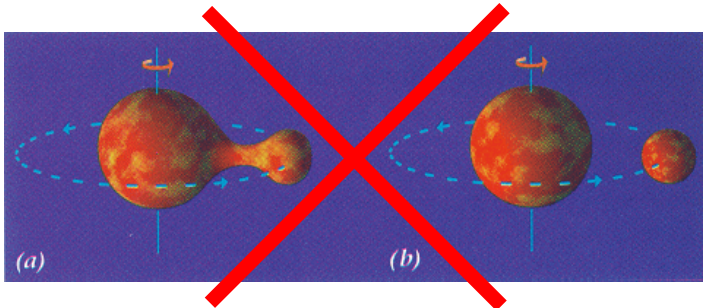
Image courtesy of Gregory H. Revera/
Wikimedia Commons

How do these constraints fit with the classical formation hypotheses?

- Can explain:
 - Lack of large core
 - Oxygen isotope similarity
- But, problems exist:
 - Earth would have had to spin very fast before fission event
 - Pacific Basin only formed 70 million years ago
 - Moon's orbit and Earth's equator not co-planar
 - Cannot account for lack of volatiles on the Moon



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- Can explain:
 - Compositional differences
- But, problems:
 - Small core
 - Isotopic similarities
 - Unlikely to be captured:
 - More likely to impact or be hurled back out to space

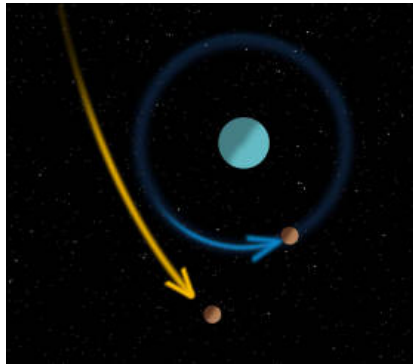


Image courtesy of Black Cat Studios

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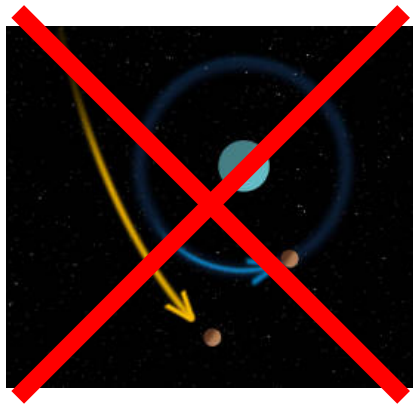
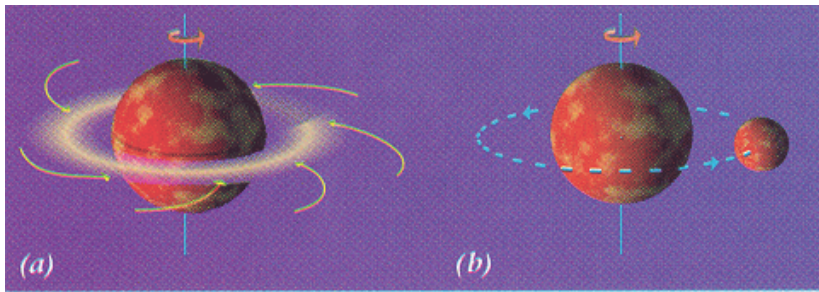


Image courtesy of Black Cat Studios

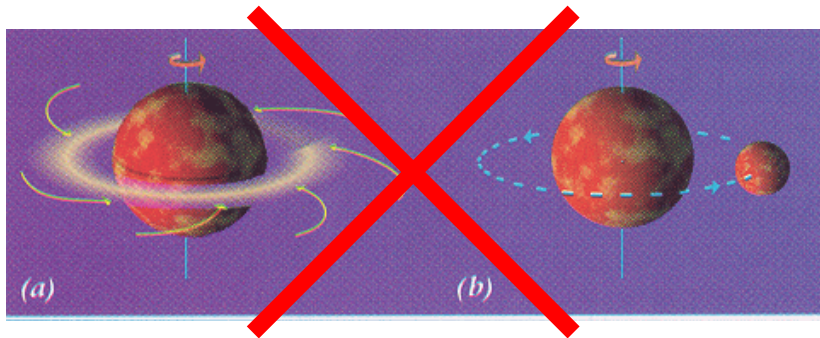
Co-accretion

- Most popular theory until 1970s
 - Doesn't require low frequency event like capture
- But, cannot explain:
 - Difference in iron content
 - Large angular momentum of Earth–Moon system

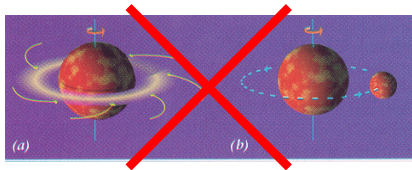
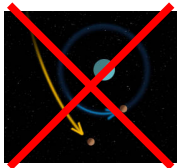
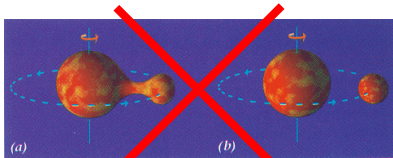


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New theory required



- Not fission, capture or co-accretion

- After Apollo, we needed a new theory to explain the origin of the Moon

Part 3:

Giant Impact Hypothesis



Image courtesy of NASA

Giant impact hypothesis

- Mid 1970's to 1980's:
New theory
- First suggested by
Hartmann & Davis
- Giant impact with
Earth
 - Debris disk
 - Coalesced to form
the Moon

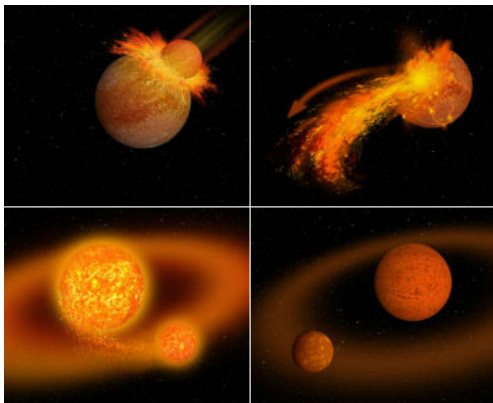
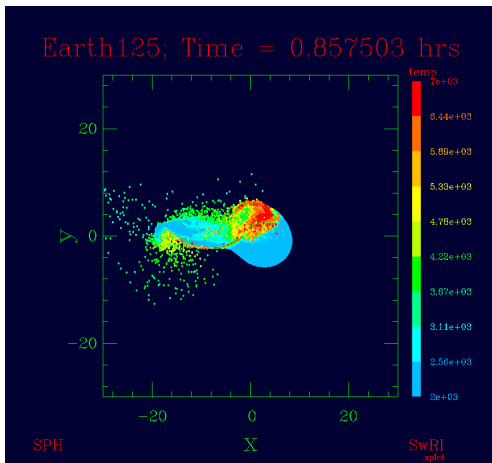


Image courtesy of Black Cat Studios

Giant impact

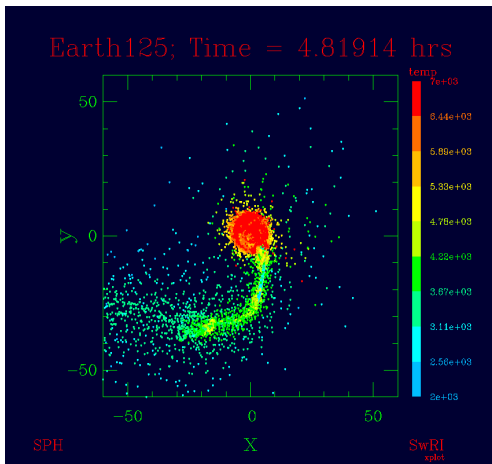
- 2004: Robin Canup's model of the giant impact
- Impact of a ~ Mars-sized body with Earth
 - **Theia**
- Impact angle 42–50°
- Moon is composed of > 60% impactor material



Images courtesy of Canup (2004) *Icarus*

Giant impact

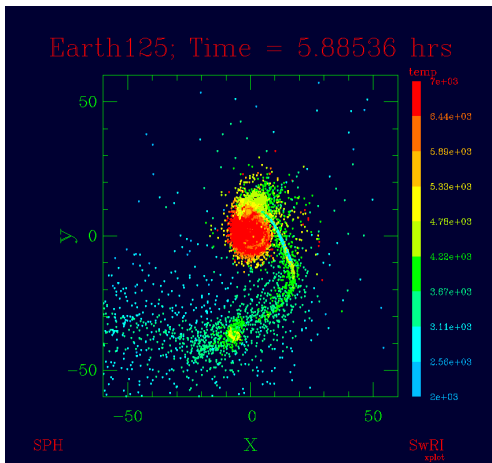
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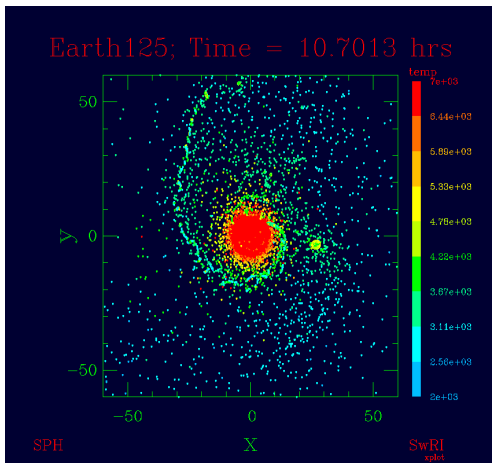
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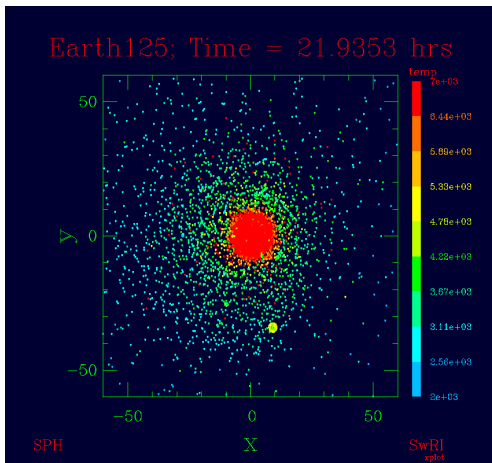
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Images courtesy of Canup (2004) *Icarus*

Giant impact

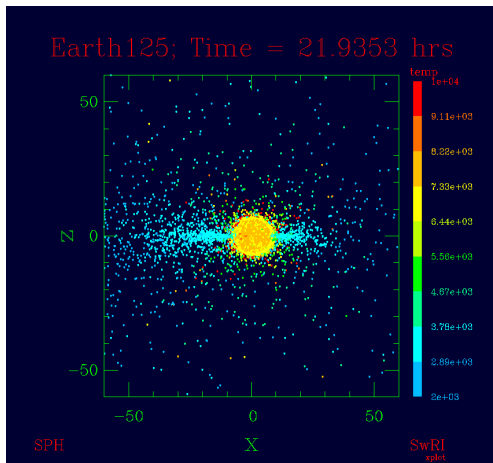
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Images courtesy of Canup (2004) *Icarus*

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In the lecture, I showed a movie of this collision. That movie, and other resources from Canup (2004), can be found at:

<http://www.boulder.swri.edu/~robin/moonimpact/>

Success?

Does this model match the observational constraints?

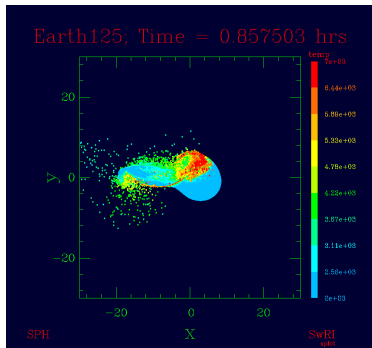


Image courtesy of Canup (2004) *Icarus*

■ Explains a lot of the constraints on Moon formation

- 1 Angular momentum ✓
- 2 Low iron content/small core ✓
- 3 Late formation time ✓
- 4 Magma ocean ✓

One constraint cannot be easily explained by this model

- Measurements of oxygen and titanium isotopes — now a lot more accurate
- As the error bars on the measurements shrunk, isotopic composition of the Moon and Earth became indistinguishable
- This means that:
 - They must come from a single body, or
 - They were well mixed during/after the impact

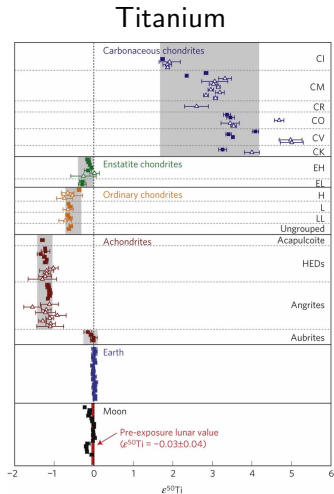


Image courtesy of Zhang et al. (2012) *Nature Geoscience*

Can we save the Giant Impact Hypothesis?

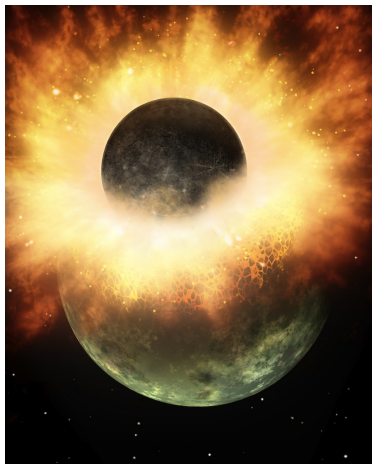


Image courtesy of NASA

- 1 Earth and Theia came from the same orbital distance
 - Explains similar isotopic composition
 - But, can't account for late formation time
- 2 Material stayed hot for a long time, allowing exchange in the vapor phase
 - Could reach isotopic equilibrium for volatile elements like oxygen
 - However, refractory elements like tungsten and titanium and silicon are not readily explained

This is where things stood, up to around a month ago

Part 4:

Giant Impact Hypothesis: *Updated*



Image courtesy of NASA

Moving the goalposts

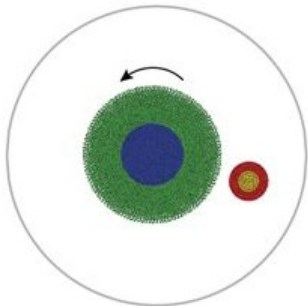


Image courtesy of Čuk & Stewart (2012)
Science

- In the original scenario, the angular momentum of the Earth–Moon system was generated by the impact
 - This limited the types of impact that could be studied
- A month ago, two papers removed this limitation
- They assumed that the Earth was spinning rapidly *before* the impact
 - A **resonance** with the Sun after the impact was able to remove the excess angular momentum

Ćuk and Stewart's rapidly spinning model

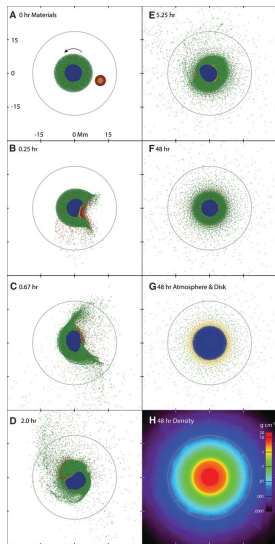


Image courtesy of Ćuk & Stewart
(2012) *Science*

- In one of the new models, a smaller impactor is invoked
 - It doesn't need to generate angular momentum
- Small mass — lower fraction of Moon's composition
- Helps explain isotopic similarities
- Think of it as a hybrid of the Giant Impact hypothesis and Darwin's Fission hypothesis

Canup's new model

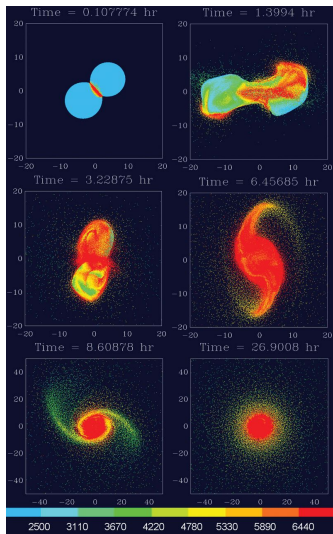


Image courtesy of Canup (2012) *Science*

- In another updated model, Canup also invokes a fast spinning Earth
- But, rather than a smaller impactor, uses an impactor more similar in size to proto-Earth
- Assumption is that the resultant bodies are a mixture of Earth and Theia
- Works well for Theia of $\sim 40\%$ the size of Earth

A hit-and-run collision

- One final alternative doesn't need to invoke resonance with the Sun to slow down the Earth's spin
- Instead, it assumes a **hit-and-run** type collision
 - High velocity, low incidence angle
- Most of the impactor mass continues downrange
- Moon forms from debris disk that is composed primarily of Earth material

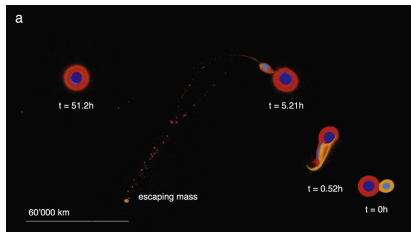


Image courtesy of Reufer et al. (2012) *Icarus*

- Impacts, once again, played a key role in shaping the Solar System
- The full details of the impact that formed our Moon are still not resolved
- Exciting work to come in the next few years
 - Which of the possible models best fits our observations?



Image courtesy of Gregory H. Revera/
Wikimedia Commons

Thank you

Questions?